

# **RothC - A model for the turnover of carbon in soil**

**Model Description**

**(February 2024)**

**K. Coleman, J. M. Prout & A. E. Milne**

**Rothamsted Research  
Harpenden  
Herts  
AL5 2JQ**

**Contact:-**

**Kevin Coleman, Jonah Prout or Alice Milne**

**E-mail:-**

kevin.coleman@rothamsted.ac.uk

jonah.prout@rothamsted.ac.uk

alice.milne@rothamsted.ac.uk

# Contents

|               |            |                                                                                                                                                  |
|---------------|------------|--------------------------------------------------------------------------------------------------------------------------------------------------|
| <b>Part 1</b> | <b>1.1</b> | <b>Introduction</b>                                                                                                                              |
|               | <b>1.2</b> | <b>Data requirements</b>                                                                                                                         |
|               | <b>1.3</b> | <b>Model structure</b>                                                                                                                           |
|               | <b>1.4</b> | <b>Decomposition of an active compartment</b>                                                                                                    |
|               | <b>1.5</b> | <b>Decomposition rate constants</b>                                                                                                              |
|               | <b>1.6</b> | <b>Calculation of the rate modifying factors</b>                                                                                                 |
|               | <b>1.7</b> | <b>Partitioning of carbon between that lost from the soil and that remaining: the <math>\text{CO}_2 / (\text{BIO} + \text{HUM})</math> ratio</b> |
|               | <b>1.8</b> | <b>Calculation of the equivalent radiocarbon age</b>                                                                                             |
|               | <b>1.9</b> | <b>Definitions of abbreviations used</b>                                                                                                         |
| <b>Part 2</b> | <b>2.1</b> | <b>Example of the use of the model</b>                                                                                                           |
|               | <b>2.2</b> | <b>How the model calculates the carbon content of the soil and its equivalent radiocarbon age</b>                                                |

## References

# Part 1

## 1.1 Introduction

RothC is a model for the turnover of organic carbon in non-waterlogged top-soils that allows for the effects of soil type, temperature, moisture content and plant cover on the turnover process. It uses a monthly time step to calculate total organic carbon ( $\text{t ha}^{-1}$ ), microbial biomass carbon ( $\text{t ha}^{-1}$ ) and  $\delta^{14}\text{C}$  (from which the equivalent radiocarbon age of the soil can be calculated) on a years to centuries timescale (Jenkinson, 1990; Jenkinson et al., 1991; Jenkinson and Coleman, 1994; Jenkinson et al., 1992; Jenkinson et al., 1987). It needs few inputs and those it needs are easily obtainable. It is an extension of the earlier model described by Jenkinson and Rayner (1977) and by Hart (1984). A comparative study of C turnover models, including RothC, was published (Smith et al., 1997). RothC is designed to run in two modes: 'forward' in which known inputs are used to calculate changes in soil organic matter and 'inverse', when inputs are calculated from known changes in soil organic matter.

RothC was originally developed and parameterized to model the turnover of organic C in arable top-soils from the Rothamsted Long Term Field Experiments - hence the name. Later, it was extended to model turnover in grassland and in woodland and to operate in different soils and under different climates. It should be used cautiously on subsoils, soils developed on recent volcanic ash (but see Saggarr et al., 1996; Shirato et al., 2004; Tate et al., 1994), soils from the tundra and taiga and not at all on soils that are permanently waterlogged. It has also been used to model changes in soil carbon in *Pinus radiata* on Mediterranean agricultural soils (Romanya et al., 2000), forest and pasture ecosystems of Amazon, Brazil (Cerri et al., 2003), Mediterranean agro-silvo-pastoral systems (Francaviglia et al., 2012).

Farina et al. (2013) modified the soil water dynamics for semi-arid regions, and Giongo et al. (2020) created a daily version and modified the soil water dynamics, for Caatinga shrublands, in the semiarid region, North-East Brazil.

## **1.2 Data requirements**

The data required to run the model are: -

- 1) Monthly rainfall (mm).
- 2) Monthly open pan evaporation (mm).

Rainfall and open-pan evaporation are used to calculate topsoil moisture deficit (TSMD), as it is easier to do this than obtain monthly measurements of the actual topsoil water deficit. If open-pan evaporation is not available, monthly potential evapotranspiration can be calculated with adequate accuracy from Müller (1982) collection of meteorological data for sites around the world. Sites should be selected from Müller's collection that are as similar climatically as possible to the site under investigation. Column 14 in Müller's Tables is headed 'Mean potential evaporation', but in fact this column gives calculated mean monthly potential evapotranspiration. If Müller's 'Mean potential transpiration' is used, you must remember to convert his values to open-pan evaporation by dividing them by 0.75. This is most important because the model is presently primed to run on open-pan evaporation data, which is then multiplied internally by 0.75 to give evapotranspiration.

i.e., if Müller's data are used as an input for the model, Open-pan evaporation = 'Mean potential transpiration' / 0.75

- 3) Average monthly mean air temperature (°C).

Air temperature is used rather than soil temperature because it is more easily obtainable for most sites. For Rothamsted, monthly air temperature satisfactorily represents the monthly mean soil temperature in topsoil, the soil temperature at 20 cm showing a difference of only +1°C of the annual minimum and -1°C of the annual maximum.

- 4) Clay content of the soil (as a percentage).

Clay content is used to calculate how much plant available water the topsoil can hold; it also affects the way organic matter decomposes.

- 5) An estimate of the decomposability of the incoming plant material - the DPM/RPM ratio.

- 6) Soil cover - Is the soil bare or vegetated in a particular month?.

It is necessary to indicate whether or not the soil is vegetated because decomposition has been found to be faster in fallow soil than in cropped soil, even when the cropped soil is not allowed to dry out (Jenkinson et al., 1987; Sommers et al., 1981; Sparling et al., 1982).

- 7) Monthly input of plant residues (t C ha<sup>-1</sup>).

The plant residue input is the amount of carbon that is put into the soil per month (t C ha<sup>-1</sup>), including carbon released from roots during crop growth. As this input is rarely known, the model is most often run in 'inverse' mode, generating input from known soil, site and weather data.

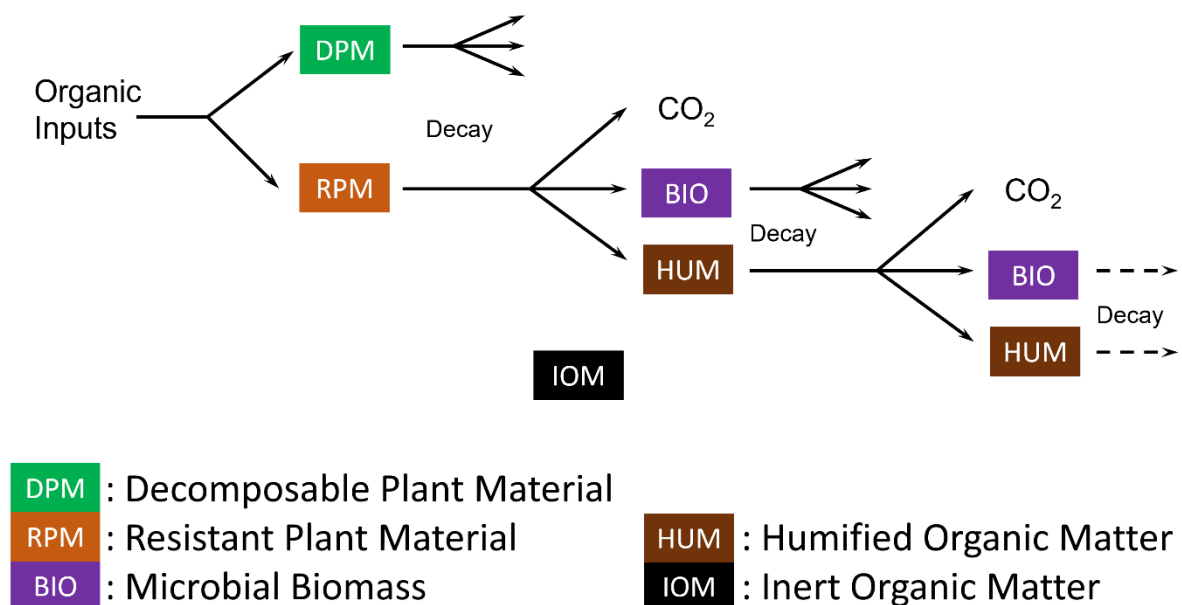
- 8) Monthly input of farmyard manure (FYM) (t C ha<sup>-1</sup>), if any.

The amount of FYM ( $\text{t C ha}^{-1}$ ) put on the soil, if any, is inputted separately, because FYM is treated slightly differently from inputs of fresh plant residues.

- 9) Depth of soil layer sampled (cm)

### **1.3 Model Structure**

Soil organic carbon is split into four active compartments and a small amount of inert organic matter (IOM). The four active compartments are Decomposable Plant Material (DPM), Resistant Plant Material (RPM), Microbial Biomass (BIO) and Humified Organic Matter (HUM). Each compartment decomposes by a first-order process with its own characteristic rate. The IOM compartment is resistant to decomposition. The structure of the model is shown in Figure 1.



**Figure 1** Structure of the Rothamsted Carbon Model.

Incoming plant carbon is split between DPM and RPM, depending on the DPM/RPM ratio of the particular incoming plant material. For most agricultural crops and improved grassland, we use a DPM/RPM ratio of 1.44 i.e., 59% of the plant material is DPM and 41% is RPM. For unimproved grassland and scrub (including Savanna) a ratio of 0.67 is used. For a deciduous or tropical woodland, a DPM/RPM ratio of 0.25 is used, so 20% is DPM and 80% is RPM. All incoming plant material passes through these two compartments once, but only once.

Both DPM and RPM decompose to form CO<sub>2</sub>, BIO and HUM. The proportion that goes to CO<sub>2</sub> and to BIO + HUM is determined by the clay content of the soil - see section 1.7. The BIO + HUM is then split into 46% BIO and 54% HUM. BIO and HUM both decompose to form more CO<sub>2</sub>, BIO and HUM.

FYM is assumed to be more decomposed than normal crop plant material. It is split in the following way: DPM 49%, RPM 49% and HUM 2%.

#### **1.4 Decomposition of an active compartment**

If an active compartment contains  $Y \text{ t C ha}^{-1}$ , this declines to  $Y e^{-abckt} \text{ t C ha}^{-1}$  at the end of the month.

where  $a$  is the rate modifying factor for temperature

$b$  is the rate modifying factor for moisture

$c$  is the soil cover rate modifying factor

$k$  is the decomposition rate constant for that compartment

$t$  is 1 / 12, since  $k$  is based on a yearly decomposition rate.

So  $Y (1 - e^{-abckt})$  is the amount of the material in a compartment that decomposes in a particular month.

#### **1.5 Decomposition rate constants**

The decomposition rate constants ( $k$ ), in years<sup>-1</sup>, for each compartment are set at: -



DPM : 10.0  
RPM : 0.3  
BIO : 0.66  
HUM : 0.02

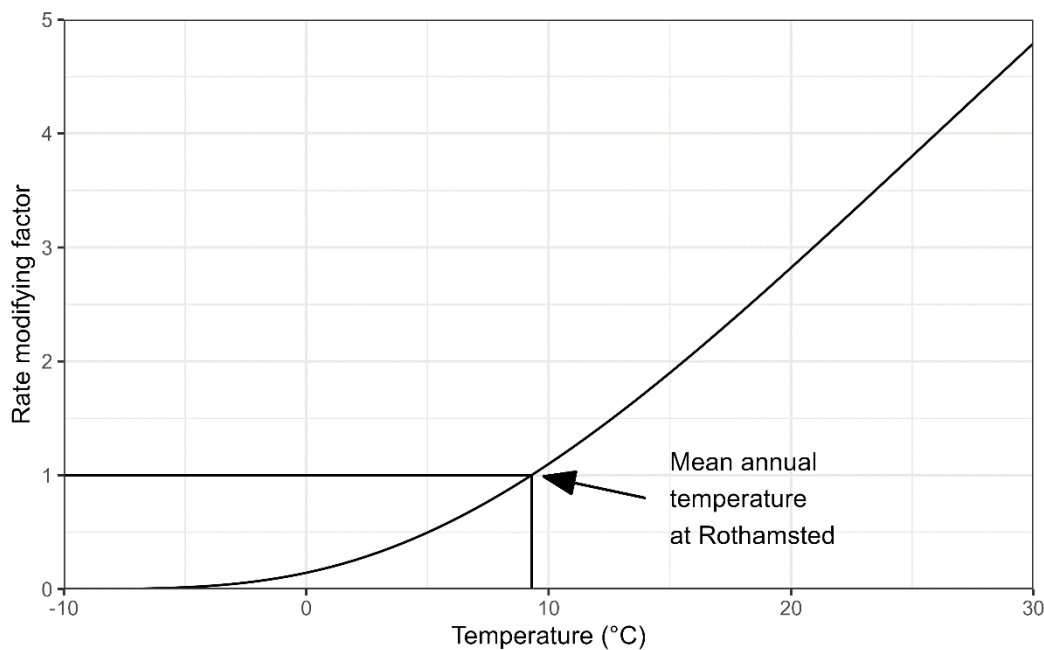
These values were originally set by tuning the model to data from some of the long-term field experiments at Rothamsted (Jenkinson et al., 1992; Jenkinson et al., 1987): they are not normally altered when using the model.

### **1.6 Calculation of the rate modifying factors**

**Temperature:** the rate modifying factor (a) for temperature is given by:

$$a = \frac{47.91}{1 + e^{\left(\frac{106.06}{T+18.27}\right)}}$$

where T is the average monthly air temperature (°C). (Figure 2)



**Figure 2** The rate modifying factor for temperature. The mean annual temperature at Rothamsted at the time of development is marked.

**Moisture:** the topsoil moisture deficit (TSMD) rate modifying factor (*b*) is calculated in the following way:-

The maximum TSMD for the 0-23 cm layer of a particular soil is first calculated from:

$$\text{Maximum TSMD} = -(20.0 + 1.3 (\% \text{clay}) - 0.01 (\% \text{clay})^2)$$

So, for Rothamsted (%clay = 23.4), the maximum TSMD = - 44.94

For a soil layer of different thickness, the maximum TSMD thus calculated is divided by 23 and multiplied by the actual thickness, in cm.

Next, the accumulated TSMD for the specified layer of soil is calculated from the first month when  $0.75 \times (\text{open pan evaporation})$  exceeds rainfall until it reaches the max. TSMD, where it stays until the rainfall starts to exceed  $0.75 \times (\text{open pan evaporation})$  and the soil wets up again. If open pan evaporation is not known, potential evapotranspiration from Müller (1982) can be used by selecting sites from his compilation that are as similar climatically to the sampling site as possible. Note that the model is presently primed to run on open-pan evaporation data, which is multiplied internally by 0.75 to give evapotranspiration. Data from Müller's tables should therefore be divided by 0.75 before entering: this is **most** important.

### **Bare soil moisture deficit (BareSMD)**

The maximum TSMD obtained above is that under actively growing vegetation: if the soil is bare during a particular month, this maximum is divided by 1.8 to give BareSMD, to allow for the reduced evaporation from a bare soil. When the soil is bare it is not allowed to dry out further than BareSMD, unless the accumulated TSMD is already less than BareSMD in which case it cannot dry out any further.

An example of this calculation for Rothamsted is shown below.

### Accumulated Topsoil Moisture Deficit (Acc. TSMD) for Rothamsted

|            | Rainfall | Open pan<br>evaporation | 0.75*E | R - 0.75*E | Acc. TSMD |
|------------|----------|-------------------------|--------|------------|-----------|
| (mm water) |          |                         |        |            |           |
| Jan        | 74       | 8                       | 6.00   | 68.00      | 0.00      |
| Feb        | 59       | 10                      | 7.50   | 51.50      | 0.00      |
| Mar        | 62       | 27                      | 20.25  | 41.75      | 0.00      |
| Apr        | 51       | 49                      | 36.75  | 14.25      | 0.00      |
| May        | 52       | 83                      | 62.25  | -10.25     | -10.25*   |
| Jun        | 57       | 99                      | 74.25  | -17.25     | -27.50    |
| Jul        | 34       | 103                     | 77.25  | -43.25     | -44.94**  |
| Aug        | 55       | 91                      | 68.25  | -13.25     | -44.94    |
| Sep        | 58       | 69                      | 51.75  | 6.25       | -38.69    |
| Oct        | 56       | 34                      | 25.50  | 30.50      | -8.19     |
| Nov        | 75       | 16                      | 12.00  | 63.00      | 0.00      |
| Dec        | 71       | 8                       | 6.00   | 65.00      | 0.00      |

\*First month when 0.75 (evaporation) is greater than the rainfall

\*\*Max. TSMD

Finally, the rate modifying factor ( $b$ ) used each month is calculated from:

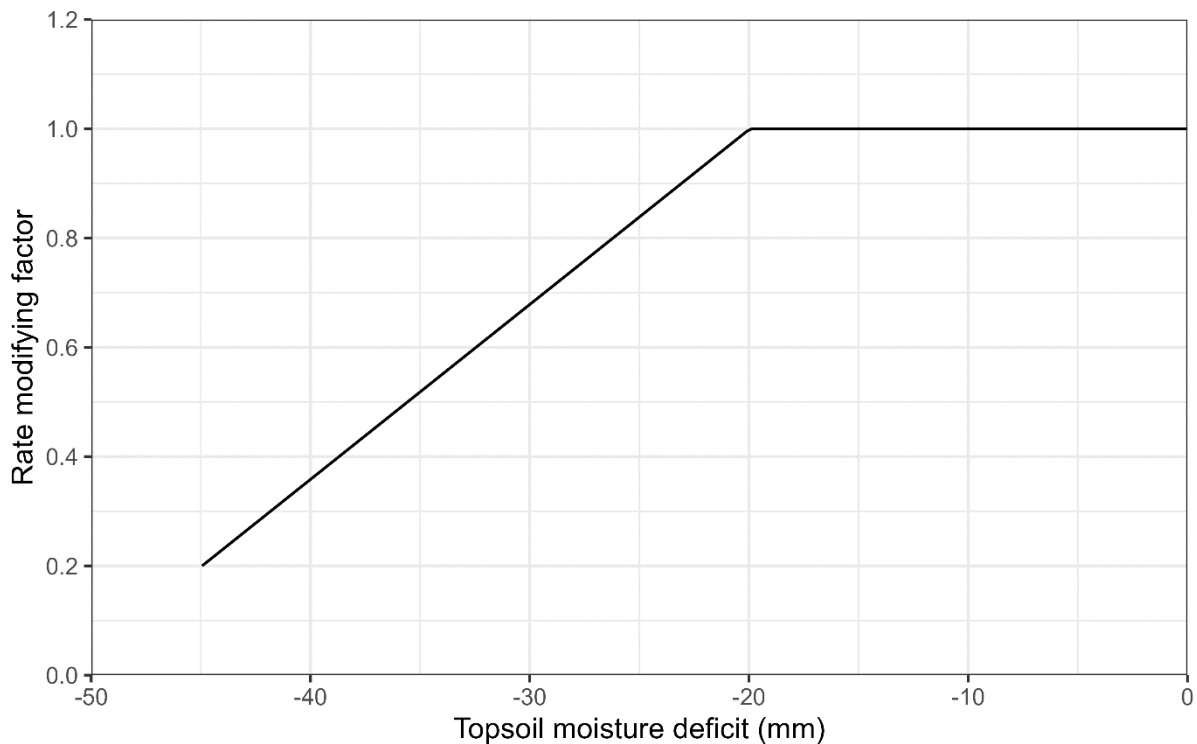
$$\text{if } acc.TSMD > 0.444 \max . TSMD,$$

$$b = 1.0$$

otherwise,

$$b = 0.2 + (1.0 - 0.2) * \frac{(\max . TSMD - acc.TSMD)}{(\max . TSMD - 0.444 \max . TSMD)}$$

This is shown in Figure 3.



**Figure 3** The rate modifying factor for moisture.

Note that the calculation in the above table starts from the 1<sup>st</sup> of January, when the soil is assumed to be at field capacity. For situations where this is not so, the weather data input should be displaced by a whole number of months, so that the soil is at field capacity at the start of the model run. Thus, in the Southern Hemisphere, the weather data file should start in July when the soil is wet, so that July will appear as January in the output.

**Soil cover factor:** The soil cover factor ( $c$ ) slows decomposition if growing plants are present. In earlier version of the model this factor is called the 'retainment factor'

If soil is vegetated  $c=0.6$

If soil is bare  $c=1.0$

### **1.7 Partitioning of carbon between that lost from the soil and that remaining: the CO<sub>2</sub> / (BIO+HUM) ratio**

The model adjusts for soil texture by altering the partitioning between CO<sub>2</sub> evolved and (BIO+HUM) formed during decomposition, rather than by using a rate modifying factor, such as that used for temperature. The ratio CO<sub>2</sub> / (BIO + HUM) is calculated from the clay content of the soil using the following equation: -

$$x = 1.67 (1.85 + 1.60 \exp(-0.0786 \% \text{clay}))$$

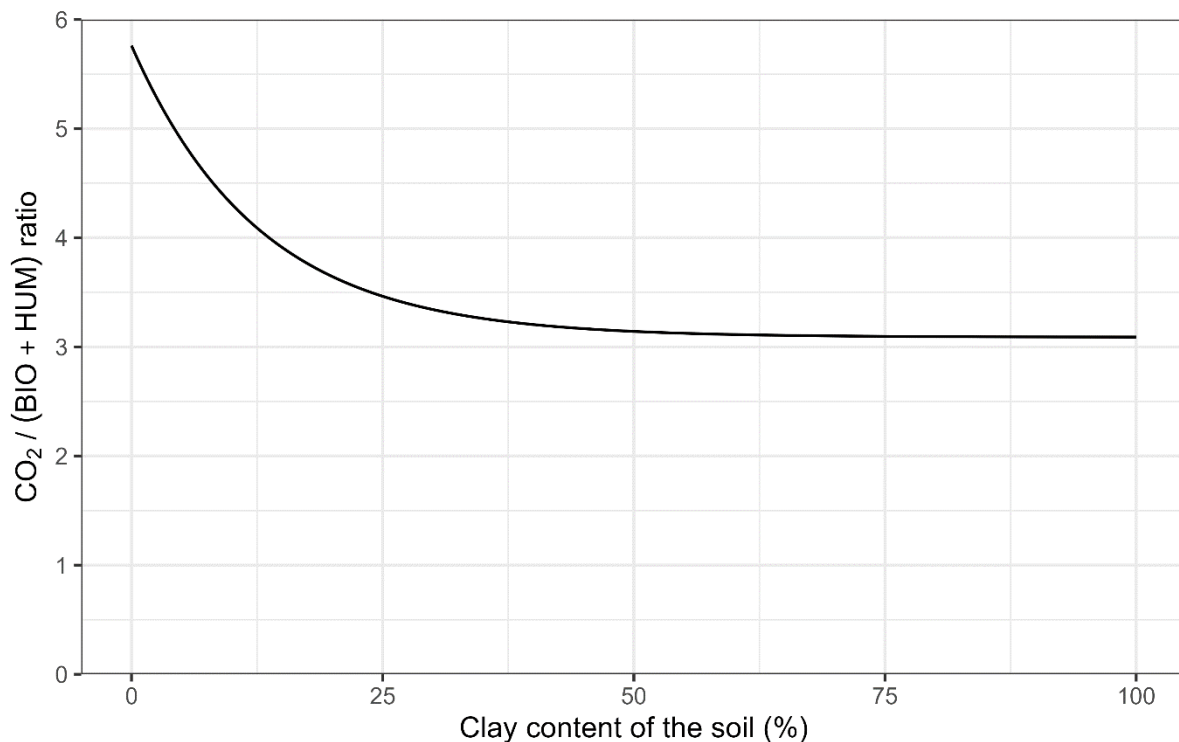
where  $x$  is the ratio CO<sub>2</sub> / (BIO+HUM)

Then  $x / (x + 1)$  is evolved as CO<sub>2</sub>

and  $1 / (x + 1)$  is formed as BIO + HUM

The scaling factor 1.67 is used to set the CO<sub>2</sub> / (BIO+HUM) ratio in Rothamsted soils (23.4% clay) to 3.51: the same scaling factor is used for all soils.

Figure 4 shows how the % clay content of the soil affects the soil texture factor, i.e. the CO<sub>2</sub> / (BIO+HUM) ratio.



**Figure 4** The effect of clay on the ratio of CO<sub>2</sub> released to (BIO+HUM) formed.

Note that the above equation relating the CO<sub>2</sub> / (BIO+HUM) ratio to %clay is not the same as that given by Jenkinson et al. (1987) or Jenkinson (1990).

### **1.8 Calculation of the equivalent radiocarbon age**

Radiocarbon measurements are commonly expressed in one of two ways, as % modern,

i.e.  $100 \text{ (specific activity of the sample) / (specific activity of the standard)}$

or as the  $\delta^{14}\text{C}$  value,

i.e.  $1000 \text{ (specific activity of the sample - specific activity of the standard) / (specific activity of the standard)}$ .

So  $\delta^{14}\text{C} = 10 \text{ (% Modern)} - 1000$

The standard is defined as 0.95 of the <sup>14</sup>C activity of the NBS standard oxalic acid.

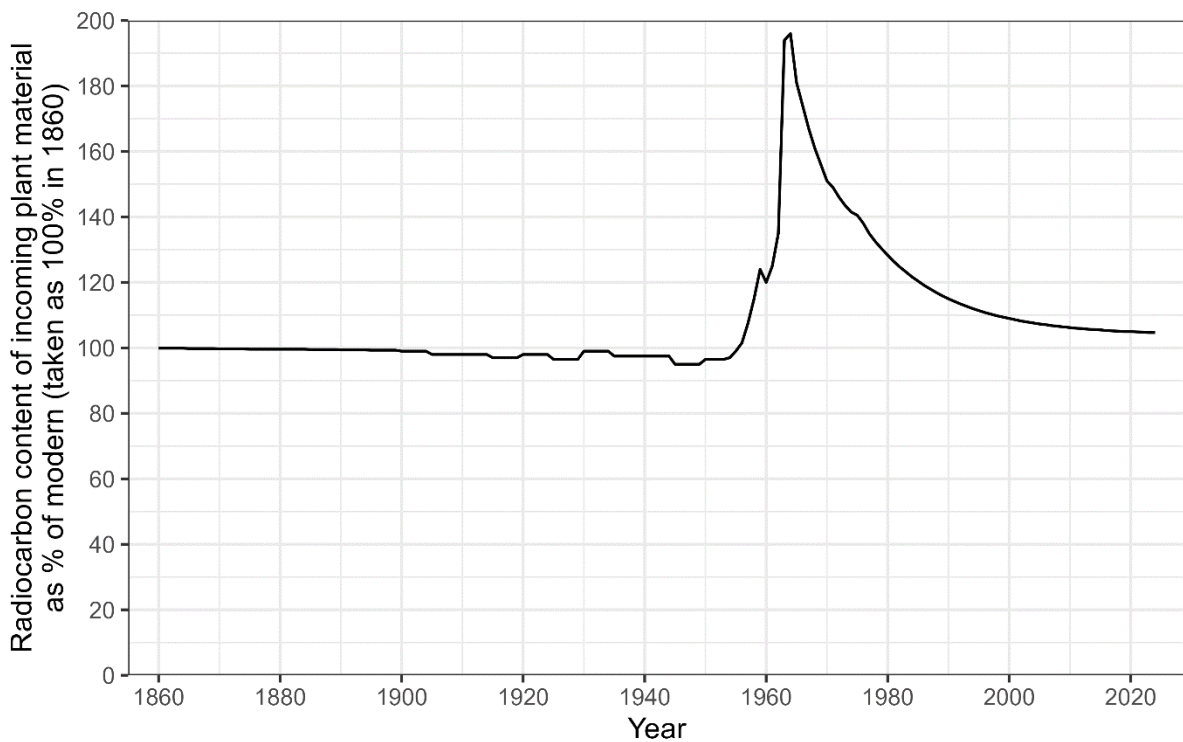
Equivalent radiocarbon age is related to  $\delta^{14}\text{C}$  in the model by the following equation

$$\delta^{14}\text{C} = 1000 \exp(-\text{equivalent radiocarbon age} / 8035) - 1000$$

using the conventional half-life for  $^{14}\text{C}$  (5568 years)

Equivalent radiocarbon age is defined as the radiocarbon age of a homogeneous sample having the same radiocarbon content as the measured (non-homogeneous) sample.

Before 1860, the model assumes that the radiocarbon age of the plant material entering the soil each year is zero, i.e. its  $\delta^{14}\text{C}$  value is zero and it is 100 % modern. After 1860 the radiocarbon content of the incoming plant carbon (expressed as % modern) in a particular year is set from an internal table - shown graphically in Figure 5.



**Figure 5** Postulated radiocarbon content of incoming plant material 1860–2024.

This input table was constructed from data on the radiocarbon content of atmospheric CO<sub>2</sub>: for the 1860 - 1949 period from Baxter and Walton (1971), for 1950 - 1984 from Harkness et al. (1986) and for 1987 - 1993 from Levin et al. (1997). The radiocarbon content of each year's input of plant carbon is taken to be the same as that of atmospheric CO<sub>2</sub> for the same year. The 'radiocarbon activity scaling factor' in the model print-out is the radiocarbon activity of the input for a particular year, expressed as either (%modern) / 100 or ( $\Delta^{14}\text{C} + 1000$ ) / 1000, i.e., taking the value for 1859 as 1.

The age of the IOM fraction is set by default to 50,000 years, implying that it contains virtually no <sup>14</sup>C ( $\Delta^{14}\text{C} = -998.0$ ) and that it is of geological age rather than pedological age.

If no radiocarbon measurements are available, IOM is set using the equation below (Falloon et al., 1998):

$$\mathbf{IOM} = 0.049 * (\mathbf{TOC})^{1.139}$$

Where, **IOM** is Inert organic matter, t C ha<sup>-1</sup> and **TOC** is Total organic carbon, t C ha<sup>-1</sup>

This is a very rough approximation for surface soils alone.



## **1.9 Definitions of abbreviations used**

|          |   |                                              |
|----------|---|----------------------------------------------|
| <i>a</i> | : | Rate modifying factor for temperature        |
| <i>b</i> | : | Rate modifying factor for moisture           |
| BIO      | : | Microbial biomass                            |
| <i>c</i> | : | Rate modifying factor for soil cover         |
| DPM      | : | Decomposable plant material                  |
| FYM      | : | Farm yard manure                             |
| HUM      | : | Humified organic matter                      |
| IOM      | : | Inert organic matter                         |
| <i>k</i> | : | Decomposition rate constant                  |
| RPM      | : | Resistant plant material                     |
| SMD      | : | Soil moisture deficit                        |
| <i>t</i> | : | time                                         |
| <i>T</i> | : | Temperature                                  |
| TSMD     | : | Top soil moisture deficit                    |
| TOC      | : | Total organic carbon (t C ha <sup>-1</sup> ) |

## Part 2

### 2.1 An example of the use of the model

The use of the model will be illustrated using data from one of the Rothamsted long-term field experiments, on the continuous cultivation of spring barley. This experiment was started on Hoosfield in 1852 and was designed to study the effects of fertilizers and FYM on the yield of barley. A detailed account can be found in the booklet: *Guide to the classical field experiments*. Rothamsted Experimental Station (Macdonald et al., 2018). None of the data from this experiment were used in setting the model parameters, so the fit obtained between model and data is an objective test of the model.

In modelling the Hoosfield data, it is first necessary to run the model to produce a starting soil organic C content that is the same as that originally present in the soil (33.8 t C ha<sup>-1</sup> in 1852, which includes 2.7 t C ha<sup>-1</sup> in IOM, as calculated from the equation of Falloon et al. (1998). Soil organic C is assumed to have been at equilibrium in 1852. The modelled plant input needed to obtain 33.8 t C ha<sup>-1</sup> in the soil is then 1.70 t C ha<sup>-1</sup> y<sup>-1</sup>. This input is distributed as follows: 0.212 t C ha<sup>-1</sup> month<sup>-1</sup> from January to July and in December, with no inputs in the other four months. This distribution is no more than a guess for the mixed arable cropping that prevailed on Hoosfield before the experiment commenced in 1852. It makes little difference to the calculated equilibrium value for total organic C or to radiocarbon age how the annual input is distributed, or even if it is all added in a single pulse. Only if the model is being used to predict annual changes in fractions with short turnover times (notably Biomass and DPM) will the input distribution appreciably affect the results. A soil cover factor of 1 was used in the months with plant inputs, zero in the other four months.

Once the starting C content has been established, land management files are created for each of the three treatments modelled in Figure 6; these are plot 7-2 (farmyard manure annually), plot 7-1 (farmyard manure annually 1852-1871, nothing thereafter) and a mean of plots 6-1 and 6-2 (both unmanured).

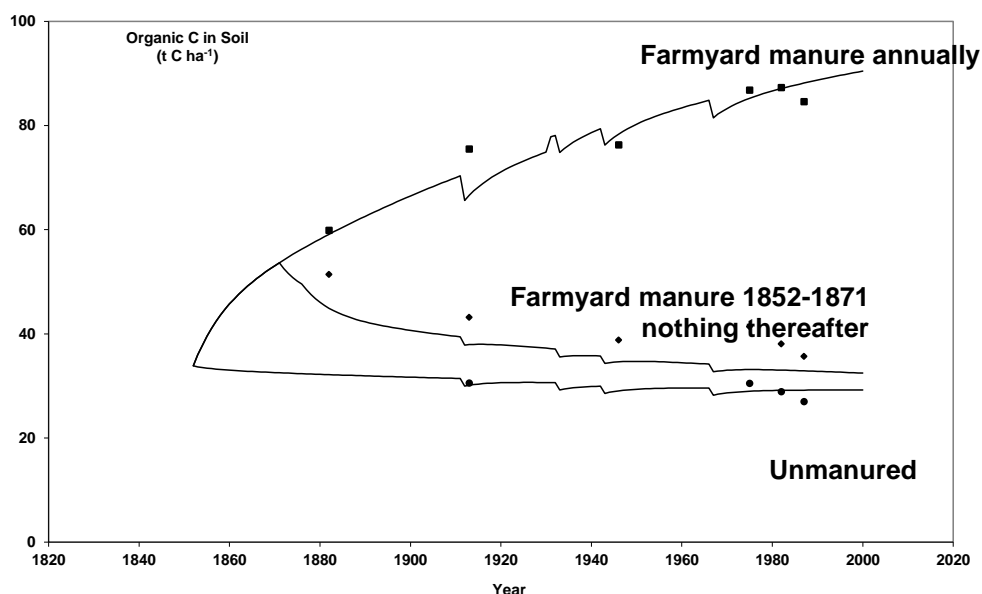
For the unmanured treatment, the annual input of plant residues was calculated to be  $1.60 \text{ t C ha}^{-1} \text{ y}^{-1}$  (distributed with  $0.16 \text{ t C ha}^{-1}$  in April,  $0.32$  in May,  $0.48$  in June and  $0.64$  in July). A soil cover factor of one was used in April, May, June and July, zero in the other months. These input figures were used from 1852 to 2000, except in the years which were fallow (1912, 1933, 1943 and 1967). For the fallow years the plant input was set at zero (bare cultivated fallow) and a soil cover factor of 0 was used in all twelve months.

For the treatment receiving farmyard manure annually (plot 7-2), the annual input of plant residues from the barley was calculated to be  $2.80 \text{ t C ha}^{-1} \text{ y}^{-1}$  ( $0.28 \text{ t C ha}^{-1}$  in April,  $0.56$  in May,  $0.84$  in June and  $1.12$  in July) Again a soil cover factor of one was used in April, May, June and July, zero in the other months. As with the unmanured treatment, this input was used from 1852 to 2000, except in the four fallow years (1912, 1933, 1943 and 1967). The FYM (containing  $3.0 \text{ t C ha}^{-1} \text{ y}^{-1}$ ) was applied in February each year from 1852-1911 and from 1913-1930. In 1931 FYM containing  $6.0 \text{ t C ha}^{-1}$  was applied ( $3.0$  in Feb and  $3.0$  in Nov). From 1932 to 2000, FYM containing  $3.0 \text{ t C ha}^{-1}$  was applied in November each year.

The third treatment received FYM ( $3 \text{ t C ha}^{-1} \text{ y}^{-1}$ ) every February from 1852 to 1871 and nothing thereafter (plot 7-1). From 1852 to 1876, plant residue input was set at  $2.80 \text{ t C ha}^{-1} \text{ y}^{-1}$  (split in the same way as plot 7-2), with the same soil cover factor of one in April, May, June and July, zero in the other months. From 1877 to 2000, plant residues were set at  $1.60 \text{ t C ha}^{-1} \text{ y}^{-1}$  (split in the same way as the unmanured plot), with the same soil cover factor, except in the four fallow years of 1912, 1933, 1943 and 1967.

Figure 6 shows the modelled data for total soil organic C in the three treatments, together with the measured data. The modelled results for the treatment receiving FYM for only 20 years are considerably lower than the measurements; agreement is closer with the other two treatments.

Figure 6 - Hoosfield continuous barley experiment  
Data modelled by RothC-26.3 (Solid lines)



## 2.2 How the model calculates the carbon content of the soil and its equivalent radiocarbon age

To see how carbon content, equivalent age and  $\Delta^{14}\text{C}$  values are calculated, we will examine the first month (January 1852) of the simulation for the unmanured plot on the Hoosfield spring barley experiment (Figure 6). First consider the position at the end of this preliminary run to equilibrium on 31<sup>st</sup> December 1851, after the model had run for 10,000 years using the Rothamsted weather file, the Hoosfield Land Management file (with an annual input of  $1.70 \text{ t C ha}^{-1} \text{ year}^{-1}$ ), a DPM/RPM ratio of 1.44 and an IOM of  $2.7 \text{ t C ha}^{-1}$ . On 31<sup>st</sup> December 1851 the state of the model is:

|       | Carbon<br>( $\text{t C ha}^{-1}$ ) | Equivalent<br>Radiocarbon<br>age (years) | $\Delta^{14}\text{C}$ |
|-------|------------------------------------|------------------------------------------|-----------------------|
| DPM   | 0.1533                             | 0.10                                     | -0.01                 |
| RPM   | 4.4852                             | 6.70                                     | -0.83                 |
| BIO   | 0.6671                             | 21.69                                    | -2.69                 |
| HUM   | 25.8576                            | 116.88                                   | -14.44                |
| IOM   | 2.7000                             | 50000.00                                 | -998.02               |
| Total | 33.8632                            | 764.37                                   | -90.75                |

Now consider the state on the 31<sup>st</sup> January 1852 for the unmanured plot, which does not receive any input of plant C or FYM in January. The temperature, moisture and soil cover during January give a combined rate modifying factor, *abc*, of 0.3561. Using the rate constants given in Section 1.5, the C content of the different compartments are changed as follows.

DPM becomes  $0.1533 * \exp[- 10 * 0.3561 / 12] = 0.1140$

RPM becomes  $4.4852 * \exp[- 0.3 * 0.3561 / 12] = 4.4455$

BIO becomes  $0.6671 * \exp[- 0.66 * 0.3561 / 12] = 0.6542$

HUM becomes  $25.8576 * \exp[- 0.02 * 0.3561 / 12] = 25.8423$

The difference between one month and the next for the C content of each compartment is:

DPM 0.0393

RPM 0.0397

BIO 0.0129

HUM 0.0153

These differences represents the material that decomposes during the month in each compartment. This material is split (see Section 1.7) between (BIO+HUM) and CO<sub>2</sub> in the following way:

$(3.51 / 4.51) * (\text{difference})$  is CO<sub>2</sub>-C

$(1 / 4.51) * (\text{difference})$  is (BIO+HUM)

The (BIO+HUM) thus formed is split as 46% BIO and 54% HUM

This is shown in the following table:

|     | Diff   | BIO    | HUM    | CO <sub>2</sub> -C |
|-----|--------|--------|--------|--------------------|
| DPM | 0.0393 | 0.0039 | 0.0047 | 0.0307             |
| RPM | 0.0397 | 0.0041 | 0.0048 | 0.0308             |
| BIO | 0.0129 | 0.0013 | 0.0015 | 0.0100             |
| HUM | 0.0153 | 0.0016 | 0.0018 | 0.0119             |

The carbon content of each compartment is now made up in the following way

$$\begin{aligned}
 \text{DPM} &= 0.1140 && = 0.1140 \\
 \text{RPM} &= 4.4455 && = 4.4455 \\
 \text{BIO} &= 0.6542 + 0.0039 + 0.0041 + 0.0013 + 0.0016 && = 0.6651 \\
 \text{HUM} &= 25.8423 + 0.0047 + 0.0048 + 0.0015 + 0.0018 && = 25.8551
 \end{aligned}$$

The model calculates the age of each compartment from a matrix which starts with the age of that compartment on 31 December 1851 and adjusts it for changes occurring during January 1852. For the DPM and RPM compartments, which in this particular example receive no fresh inputs of plant material in January, the age on 31 December is increased by one month to give the age on 31 January. For the BIO and HUM compartments, the incoming material added at the end of the month comes tagged with the age of the compartment from which it came. The age of the whole compartment is then obtained by weighting the age of its components by their carbon content. The resulting values for equivalent radiocarbon age and  $\Delta^{14}\text{C}$  are then:

|     | Equivalent<br>Radiocarbon<br>age (years) | $\Delta^{14}\text{C}$ |
|-----|------------------------------------------|-----------------------|
| DPM | 0.19                                     | -0.02                 |
| RPM | 6.78                                     | -0.84                 |
| BIO | 21.78                                    | -2.70                 |
| HUM | 116.91                                   | -14.45                |

A similar procedure is followed if there is an input of fresh plant residue during a particular month. This input is given the appropriate radiocarbon scaling factor for the year in which it occurs and distributed between DPM and RPM in the specified proportions at the end of the month in question.

For the unmanured plot in the Hoosfield experiment the calculated radiocarbon age of the whole soil organic C is 987 years in 1950 and 70 years in 1970, the decline being due to radiocarbon from thermonuclear testing. No measurements of radiocarbon are available from the Hoosfield experiment: had they been, the IOM content of the soil *and* the annual inputs of plant C would have been iteratively adjusted to give both the correct organic C content and the correct radiocarbon content for a particular sampling date.

## References

- Baxter MS, Walton A. Fluctuations of atmospheric carbon-14 concentrations during past century. *Proceedings of the Royal Society of London Series a-Mathematical and Physical Sciences* 1971; 321: 105-127.
- Cerri CEP, Coleman K, Jenkinson DS, Bernoux M, Victoria R, Cerri CC. Modeling soil carbon from forest and pasture ecosystems of Amazon, Brazil. *Soil Science Society of America Journal* 2003; 67: 1879-1887.
- Falloon P, Smith P, Coleman K, Marshall S. Estimating the size of the inert organic matter pool from total soil organic carbon content for use in the Rothamsted carbon model. *Soil Biology & Biochemistry* 1998; 30: 1207-1211.
- Farina R, Coleman K, Whitmore AP. Modification of the RothC model for simulations of soil organic C dynamics in dryland regions. *Geoderma* 2013; 200: 18-30.
- Francaviglia R, Coleman K, Whitmore AP, Doro L, Urracci G, Rubino M, Ledda L. Changes in soil organic carbon and climate change - Application of the RothC model in agro-silvo-pastoral Mediterranean systems. *Agricultural Systems* 2012; 112: 48-54.
- Giongo V, Coleman K, Santana MD, Salviano AM, Olszveski N, Silva DJ, et al. Optimizing multifunctional agroecosystems in irrigated dryland agriculture to restore soil carbon - Experiments and modelling. *Science of the Total Environment* 2020; 725.
- Harkness DD, Harrison AF, Bacon PJ. The temporal distribution of bomb carbon-14 in a forest soil. *Radiocarbon* 1986; 28: 328-337.
- Hart PBS. Effects of soil type and past cropping on the nitrogen supplying ability of arable soils. University of Reading, UK, 1984.
- Jenkinson DS. The Turnover of Organic-Carbon and Nitrogen in Soil. *Philosophical Transactions of the Royal Society of London, Series B: Biological Sciences* 1990; 329: 361-368.
- Jenkinson DS, Adams DE, Wild A. Model Estimates of Co<sub>2</sub> Emissions from Soil in Response to Global Warming. *Nature* 1991; 351: 304-306.
- Jenkinson DS, Coleman K. Calculating the Annual Input of Organic-Matter to Soil from Measurements of Total Organic-Carbon and Radiocarbon. *European Journal of Soil Science* 1994; 45: 167-174.
- Jenkinson DS, Harkness DD, Vance ED, Adams DE, Harrison AF. Calculating Net Primary Production and Annual Input of Organic-Matter to Soil from the Amount and Radiocarbon Content of Soil Organic-Matter. *Soil Biology & Biochemistry* 1992; 24: 295-308.
- Jenkinson DS, Hart PBS, Rayner JH, Parry LC. Modelling the turnover of organic matter in long-term experiments at Rothamsted. *INTECOL Bulletin* 1987; 15: 1-8.
- Jenkinson DS, Rayner JH. Turnover of soil organic matter in some of the Rothamsted classical experiments. *Soil Science* 1977; 123: 298-305.
- Levin I, Kromer B, Schoch-Fischer H, Levin I, Bruns M, Munnich M, et al. Delta 14 CO<sub>2</sub> records from two sites in Central Europe. CDIAC, ESS-DIVE repository. 1997;
- Macdonald A, Poulton P, Clark I, Scott T, Glendining M, Perryman S, et al. Guide to the Classical and Other Long-term experiments, Datasets and Sample Archive. Rothamsted Research", Rothamsted Research, 2018, pp. 57.
- Müller MJ. Selected climatic data for a global set of standard stations for vegetation science: Dr. W. Junk, Publishers. The Hague, 1982.
- Romanya J, Cortina J, Falloon P, Coleman K, Smith P. Modelling changes in soil organic matter after planting fast-growing *Pinus radiata* on Mediterranean agricultural soils.

- European Journal of Soil Science 2000; 51: 627-641.
- Saggar S, Parshotam A, Sparling GP, Feltham CW, Hart PBS. C-14-labelled ryegrass turnover and residence times in soils varying in clay content and mineralogy. *Soil Biology & Biochemistry* 1996; 28: 1677-1686.
- Shirato Y, Hakamata T, Taniyama I. Modified Rothamsted carbon model for andosols and its validation: Changing humus decomposition rate constant with pyrophosphate-extractable Al. *Soil Science and Plant Nutrition* 2004; 50: 149-158.
- Smith P, Smith JU, Powlson DS, McGill WB, Arah JRM, Chertov OG, et al. A comparison of the performance of nine soil organic matter models using datasets from seven long-term experiments. *Geoderma* 1997; 81: 153-225.
- Sommers LE, Gilmour CM, Wildung RE, Beck SM. The Effect of Water Potential on Decomposition Processes in Soils. *Water Potential Relations in Soil Microbiology*, 1981, pp. 97-117.
- Sparling GP, Cheshire MV, Mundie CM. Effect of barley plants on the decomposition of c14-labeled soil organic-matter. *Journal of Soil Science* 1982; 33: 89-100.
- Tate KR, Giltrap DJ, Parshotam A, Hewitt AE, Ross DJ, Kenny GJ, Warrick RA. Impacts of climate change on soils and land systems in New Zealand. In: Bouma WJ, Pearman GI, Manning MR, editors. *C S I R O*, Wellington, New Zealand, 1994, pp. 190-204.